

METEOROLOGY

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Lecture 2 –RADIATION

2.1 EARTH-SUN GEOMETRY

- The Earth has an elliptical orbit around the sun
- The average Earth-Sun distance is **93,000,000** mi (**150,000,000** km)
- ***Aphelion*** – the farthest Earth-Sun distance occurs in July, and is about **94,000,000** mi (**152,000,000** km)
- ***Perihelion*** – the closest Earth-Sun distance occurs in January, and is about **91,000,000** mi (**147,000,000** km)
- Astronomical seasons and climatological seasons
 - *Astronomical seasons* are :-
 - ***Winter*** – Winter Solstice to Spring Equinox (Dec **21** – Mar **21**)
 - ***Spring*** – Spring Equinox to Summer Solstice (Mar **21**- Jun **21**)
 - ***Summer*** – Summer Solstice to Fall Equinox (Jun **21** – Sep **21**)
 - ***Autumn*** – Fall Equinox to Winter Solstice (Sep **21** – Dec **21**)
 - *Climatological seasons* :-
 - ***Winter*** – December, January, February
 - ***Spring*** – March, April, May
 - ***Summer*** – June, July, August
 - ***Autumn*** – September, October, November

2.2 ENERGY, HEAT, AND TEMPERATURE

- *Energy is defined as the ability to do work*
- Types of energy
 - *Kinetic energy – energy of motion*
 - *Potential energy – stored energy*
- *Thermal Energy is defined as the total kinetic energy of all the atoms and molecules that make up a substance*
- *Temperature is a measure of the average of kinetic energy of all the atoms or molecules of the substance*
- Adding heat to a substance raises its temperature
- Subtracting heat will lower its temperature
- Methods of energy transfer
 - Conduction
 - Convection
 - Meteorologists refer to horizontal convection as *advection*
 - Radiation

2.3 RADIATION

- Radiation is comprised of electromagnetic waves of differing wavelengths as shown in table (2.1)
- Shorter wavelengths (higher frequencies) are more energetic
- Relationship between wave speed (c), wavelength (λ), and frequency (ν)

$$c = \lambda \nu$$

- Radiation moves at a speed of $c = 3.0 \times 10^8$ m/s in a vacuum
- Travels slower through matter
- Index of refraction (n) of a medium is speed in a vacuum (c) divided by speed in the medium (v)

$$n = \frac{c}{v}$$






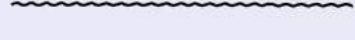

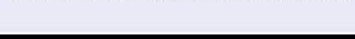
| TYPE OF RADIATION | RELATIVE WAVELENGTH | TYPICAL WAVELENGTH (meters) | ENERGY CARRIED PER WAVE OR PHOTON |
|-------------------|---|-----------------------------|---|
| AM radio waves |  | 100 | <div>Increasing</div>  |
| Television waves |  | 1 | |
| Microwaves |  | 10^{-3} | |
| Infrared waves |  | 10^{-6} | |
| Visible light |  | 5×10^{-7} | |
| Ultraviolet waves |  | 10^{-7} | |
| X rays |  | 10^{-9} | |

Table (2.1) : Types of radiation

- Electromagnetic spectrum as shown in figure (2.1)
- Visible light comprises only a small part of the spectrum (**0.4 – 0.7 μm**)
- Ultra-violet (**UV**) radiation has shorter wavelengths than visible ($\lambda < \mathbf{0.4 \mu\text{m}}$)
- Infrared (**IR**) radiation has longer wavelengths than visible ($\lambda > \mathbf{0.7 \mu\text{m}}$)
- Short wave radiation $\lambda < \mathbf{4 \mu\text{m}}$
- Long wave radiation $\lambda > \mathbf{4 \mu\text{m}}$

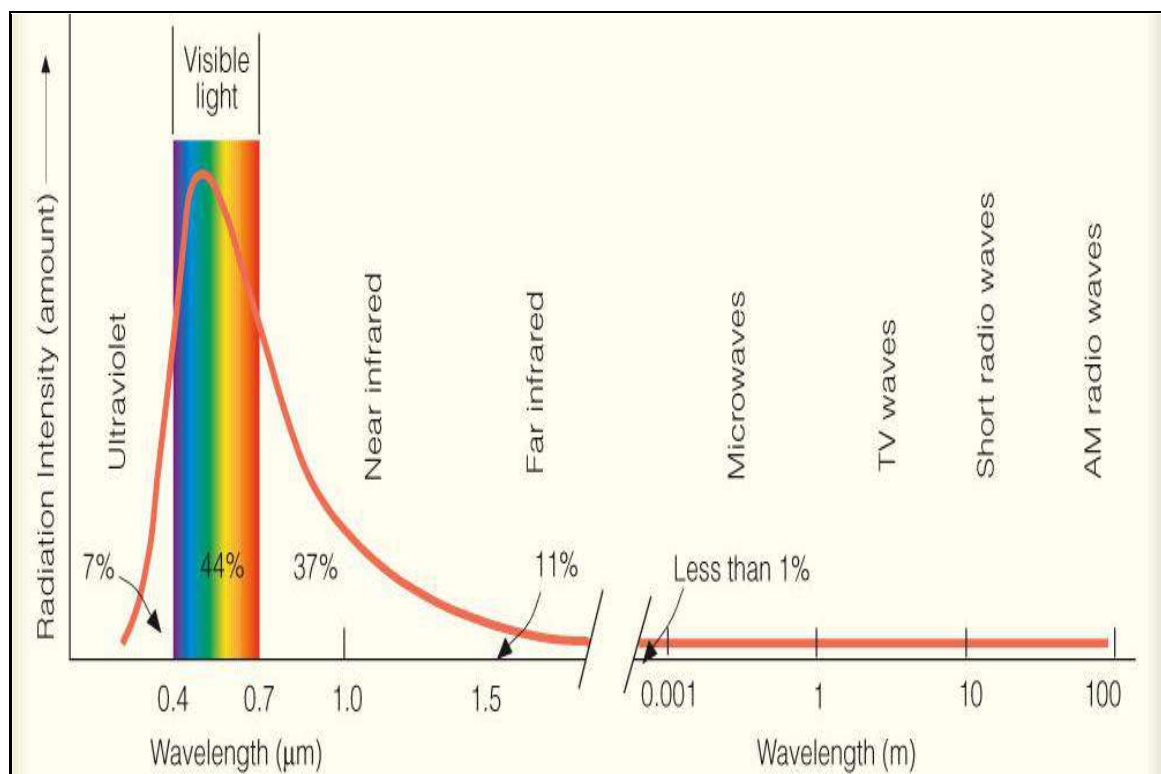


Fig. (2.1) : The sun electromagnetic spectrum

2.4 RADIATION LAWS

○ ***Planck's Law***

- It has been determined experimentally that the intensity of radiation emitted by a blackbody
- Blackbody , a hypothetical object that absorbs all of the radiation that strikes it , it also emits radiation at a maximum rate for its given temperature
- The shorter the wavelength the greater the energy

$$E_{\lambda} = \frac{c_1}{\lambda^5 \cdot (e^{c_2 / \lambda T} - 1)}$$

Where :

$$c_1 = 3.74 \times 10^{-16} \text{ W m}^2; c_2 = 1.44 \times 10^{-2} \text{ m K}, E_{\lambda} \text{ has units of } \text{W m}^{-2} \mu\text{m}^{-1}$$

○ ***Stefan-Boltzmann Law***

- A law of radiation which **states** that the amount of radiant energy emitted from a unit surface area of an object (ideally a blackbody) is proportional to the fourth power of the object's absolute temperature
- Hotter objects radiate more energy per unit area than do cold objects

$$E = \sigma T^4$$

Where :

$$\sigma = 5.67 \times 10^{-8} \text{ W m}^{-2} \text{ K}^{-4}$$

E : has units of W m^{-2}

- **Wien's Displacement Law**

- A law of radiation which **states** that the wavelength of maximum emitted radiation by an object (ideally a blackbody) is inversely proportional to the object's absolute temperature
- The hotter the object the shorter the wavelength of maximum radiation

$$\lambda_{\max} = \frac{2897 \mu m.K}{T}$$

λ_{\max} : the wavelength of maximum radiation

2.5 SOLAR (SHORTWAVE) RADIATION

- The Sun emits at about **6000 K**, while the earth emits at about **288 K** as shown in figure (2.2)
- **From Wien's displacement law**
 - The Sun emits its peak radiation at a wavelength of **0.48μm**, which is in the visible part of the spectrum (blue light)
 - The Earth emits its peak radiation at a wavelength of **10μm**, which is in the IR part of the spectrum

- **From the Stefan-Boltzman law**
 - The Sun emits $7.35 \times 10^7 \text{ W m}^{-2}$ of energy per unit area
 - The Earth emits 390 W m^{-2} of energy per unit area

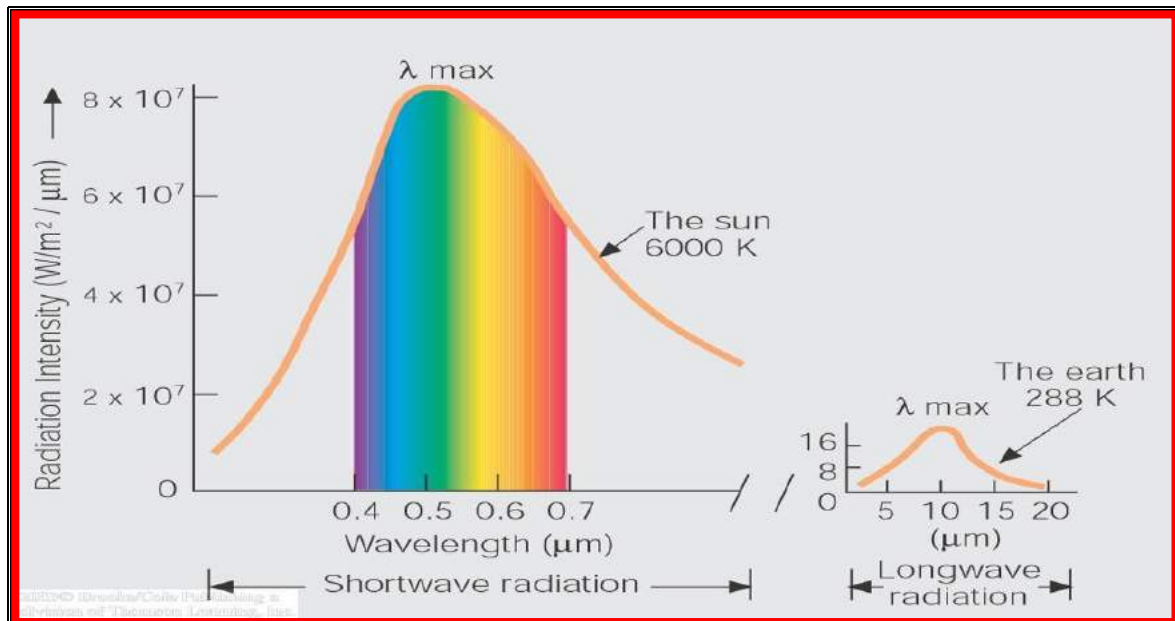


Fig. (2.2) : Short waves and Long waves Radiation

- Incoming solar radiation can be absorbed, reflected, or scattered by the atmosphere or Earth's surface as shown in figure (2.3)
 - **Absorption**
 - **19%** absorbed directly in the atmosphere, primarily by ozone in the **stratosphere**, oxygen and nitrogen in the **thermosphere**, and water vapor and clouds in the **troposphere**
 - **51%** absorbed by land and sea
 - **Reflection**
 - **30%** reflected back to space by clouds (**20%**) , the atmosphere (**6%**) and the Earth's surface (**4%**)

- *Albedo* is defined as the fraction of radiation reflected by a surface
- The Albedo of the Earth is **0.3**
- **Scattering**
 - Produces a bunch of weaker rays that move off in different directions
 - Scattering from small molecules is dependent of wavelength, with short wavelengths scattered more than longer wavelengths
 - Scattering off of larger particles is not wavelength dependent

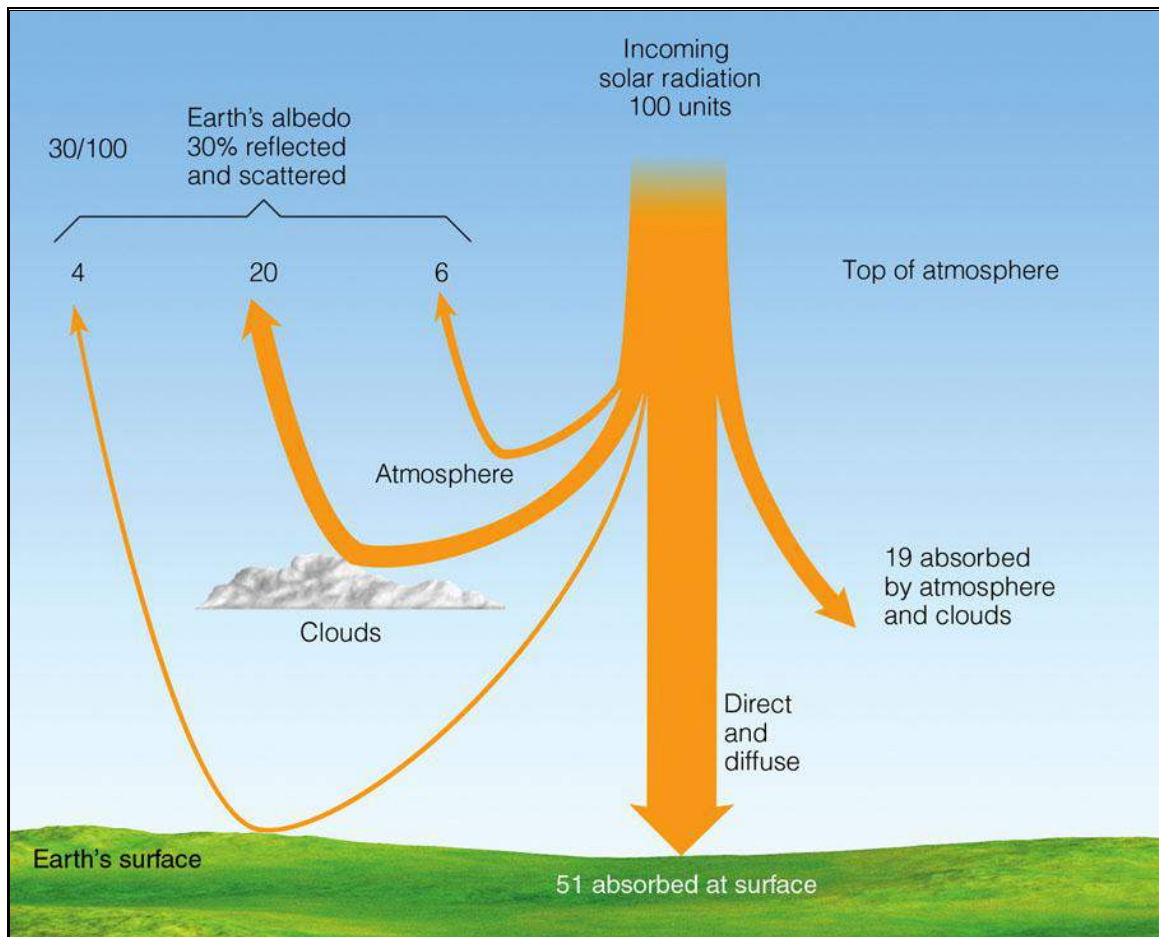


Fig. (2.3) : Fate of incoming Solar Radiation (Solae energy budget)

2.6 TERRESTRIAL (LONGWAVE) RADIATION

- Earth's surface emits radiation at a temperature of **288 K**
- Maximum emission is in the near infrared
- Almost all emission is between **5** and **25 μm**
- The atmosphere absorbs and reradiates much of the emitted **IR** radiation called "**greenhouse effect**."
- The "**greenhouse**" gases:-
 - H_2O
 - CO_2
 - N_2O
 - CH_4
 - O_3
- **Clouds** have a dual role in the radiation balance of the earth
 - Reflect incoming solar radiation, keeping daytime temperatures lower
 - Absorb outgoing terrestrial radiation, keeping nighttime temperatures higher
- Latitudinal heat balance
 - Tropics have net surplus of heat
 - Polar Regions have net deficit of heat
 - Heat must be transported from tropics to poles

2.7 RADIATION BALANCE

- In order to remain at a constant temperature, an object (earth-atmosphere) must emit the same amount of energy that it receives as shown in figure (2.4)
- The amount of solar energy absorbed by the earth and its atmosphere is

$$E_{solar} = \pi R^2 . S . (1 - \alpha) \dots (1)$$

- S is the radiation flux at the top of the atmosphere, and is called the *solar constant*, the rate at which solar energy is received on a surface at the outer edge of the atmosphere perpendicular to the sun's rays when the earth is at a mean distance from the sun
- The value of the solar constant is about **1368 W/m²**
- α is the albedo

- The amount of energy radiated by the earth and its atmosphere is

$$E_{earth} = 4\pi R^2 \cdot \sigma T^4 \dots (2)$$

- In order to be in equilibrium, E_{solar} must equal E_{earth} , so equating (1) and (2) we get

$$\pi R^2 \cdot S \cdot (1 - \alpha) = 4\pi R^2 \cdot \sigma T^4 \dots (3)$$

which can be solved for T to get , the radiation temperature of earth

$$T = \sqrt[4]{\frac{S \cdot (1 - \alpha)}{4\sigma}} \dots (4)$$

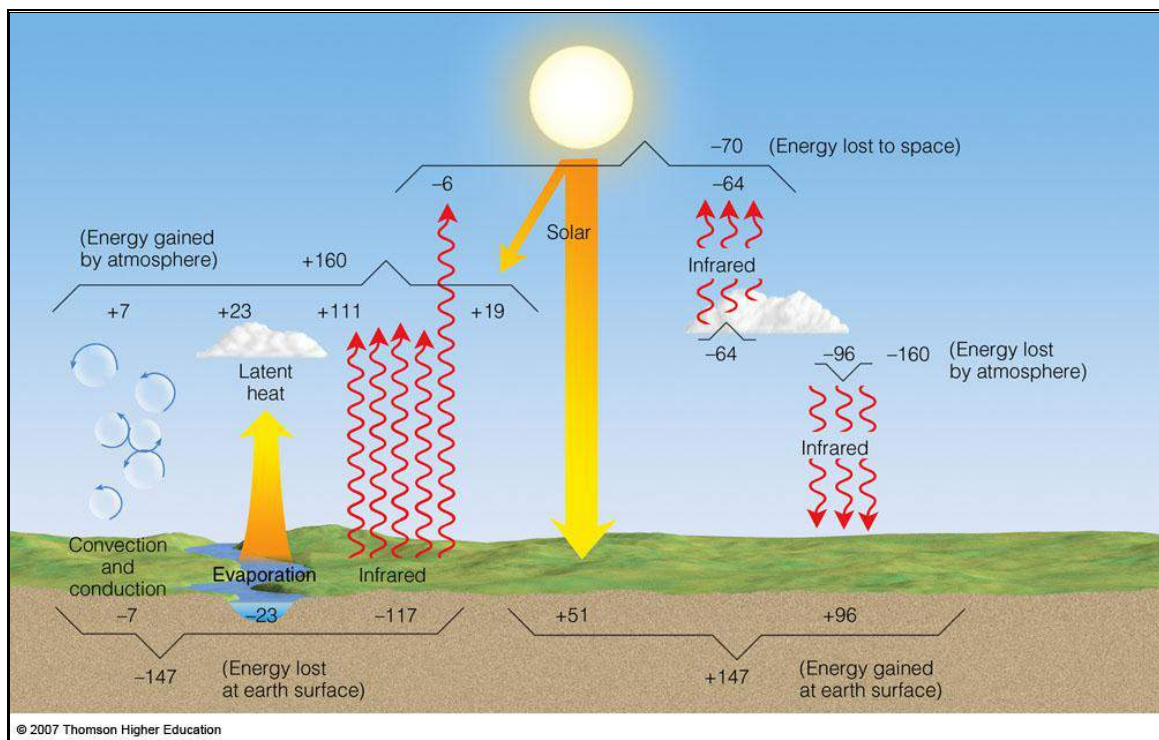


Fig. (2.4) : The earth – atmosphere energy balance